



HAL
open science

Slingram EMI Devices for Characterizing Resistive Features Using Apparent Conductivity Measurements: check of the DualEM-421S Instrument and Field Tests

Michel Dabas, Antoine Anest, Julien Thiesson, Alain Tabbagh

► **To cite this version:**

Michel Dabas, Antoine Anest, Julien Thiesson, Alain Tabbagh. Slingram EMI Devices for Characterizing Resistive Features Using Apparent Conductivity Measurements: check of the DualEM-421S Instrument and Field Tests. *Archaeological Prospection*, 2016, 23 (3), pp.165-180. 10.1002/arp.1535 . halshs-01513140

HAL Id: halshs-01513140

<https://shs.hal.science/halshs-01513140v1>

Submitted on 25 Apr 2017

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

1 **Slingram EMI devices for characterizing resistive features using apparent conductivity**
2 **measurements: check of the DualEM-421S instrument and field tests.**

3

4 Michel Dabas¹, Antoine Anest², Julien Thiesson², Alain Tabbagh²

5

6 ¹GEOCARTA, 5 rue de la banque, F-75002 ; Paris

7 ²Sorbonne Université, UPMC-Paris 6, UMR7619-Métis, F-75252, Paris

8

9 **Abstract**

10 This paper addresses the characterization of resistive archaeological targets and near surface
11 structures by ElectroMagnetic Induction (EMI). It presents tests achieved with the
12 DualEM421S instrument (Duaem Inc., Canada) in order to be able to quantitatively compare
13 these measurements to the standard technique of direct-current (DC) resistivity. The test was
14 done over the Gallo-roman site of Vieil-Evreux in Normandy, France and 1D and 3D
15 inversions were applied to the data set obtained.

16 We have first investigated the signal/noise ratio of each of the 6 DualEM receiver coils both
17 in a static mode and for a quad-pulled system. The dependence on the roll angle was also
18 measured and it is shown that rotation of DualEM must be taken into account if the roll angle
19 is more than $\pm 10^\circ$. Absolute calibration and in-phase/quadrature (out of phase) component
20 discrimination was checked by measuring the response of a small conductive and non-
21 magnetic sphere. Several EM soundings by measuring the instrument response at different
22 heights were done in order to check the quadrature (out-of-phase) response of the instrument.
23 Inversions of these EM soundings were compared to DC Vertical Electric Soundings (VESs)
24 over 4 locations and found in accordance. Several maps using different coil configurations
25 (HCP, VCP, PERP) and different heights were performed and inverted, both for a wide mesh

26 (5m) and for a finer one (0.5m). The wide mesh allows a global and rapid description of the
27 surface geology context (continuous DC measurements cannot deliver equivalent depth of
28 investigation). The fine mesh conductivity maps clearly show the walls of a *fanum* (temple) as
29 well as other structures and prove that the DualEM-421S is able to map correctly
30 archaeological resistive targets. These maps and their interpretations were compared to
31 previous results obtained by DC technique.

32

33 **Keywords**

34 EMI, multi-receivers, DualEM-421S, calibration, conductivity response, inversion,

35

36

37 **Introduction**

38 In archaeology, magnetometry is the primary technique for the detection of low
39 resistivity earthen targets, like ditches for example, while Direct Current, DC, (earth-
40 resistance) and, more recently, Ground Penetrating Radar, GPR, predominate in the detection
41 of resistive targets, like walls for example. Since both DC and GPR are expensive to acquire
42 and/or process, the detection of resistive features is likely underestimated in most of the
43 assessment of archaeological potential despite the high occurrence of resistive targets in many
44 of the archaeological sites.

45 The use of earth-resistance in the field has two main drawbacks compared to
46 magnetometry: it is more time consuming despite the development of quad-pulled systems
47 such as ARP© (Dabas, 2009), and it might not be practical to use in areas of high contact
48 resistance. The two main advantages of DC resistivity compared to magnetometry are the
49 strong response of both earthen and stone built archaeological features due to a generally high
50 electrical contrast, and also the possibility to achieve more accurate inversions using several

51 depths of investigation. Globally, despite the development of electrostatic multipole arrays
52 (Flageul *et al.* 2013) that fully overcame the galvanic contact limitation, the deployment of
53 such systems remains challenging due to their weight and complexity.

54 EM instrumentation tends to be lighter and thus more convenient than DC
55 instrumentation. These characteristics have made EMI a promising alternative to DC, but only
56 recently have instruments with an accurate phase separation and with multi-receivers become
57 readily available. These developments should open the possibility to have a conductivity
58 inversion and also to improve the use of the in-phase response. A first study aiming at a
59 global comparison of available instruments was done in the area of soil mapping in 2009
60 (Gebbers *et al.*, 2009). In the archaeology field, only a few publications have begun to tackle
61 the use of multi-receivers instruments (Saey *et al.*, 2012; Bonsall *et al.*, 2013) and none to our
62 knowledge have studied the response of multi-receivers systems to resistive targets and
63 compared EMI measurements with resistivity over a whole site.

64 Numerous archaeological sites exhibit stone built remains or compacted horizons that
65 constitute the essence of the site and provide the key to its overall organization and evolution
66 through time. These features, in general, correspond to a marked contrast in electrical
67 resistivity with the surrounding soil and the DC resistivity method has been, and is, logically
68 applied to map apparent resistivity above them. Very good results have been obtained in
69 different climatic and chrono-cultural contexts (Papadopoulos *et al.* 2009, Bossuet *et al.* 2012)
70 and the application of multi-depth arrays allowed exploring the depth and the thickness of the
71 different features (Brinon *et al.* 2012). The sensitivity of EMI devices to resistive targets is
72 indeed worse than that of the DC resistivity technique (Tabbakh 1986(a)) but these devices
73 allow simultaneous measurement of the magnetic susceptibility. Susceptibility can show for
74 example contrasts between stones and their surroundings (Gaffney *et al.* 2000, Saey *et al.*
75 2013), and experiments have shown the possibility of wall detection even with small coil

76 separation instruments (Thiesson *et al.* 2009). Compared to DC deployments for a given depth
77 of exploration, EMI instruments are more compact, which facilitates deeper investigations.

78 In this paper, after a field test in which noise and calibration of DualEM-421S were
79 analyzed, we examine the results obtained with this instrument over some buried Gallo-roman
80 walls at the site of Vieil-Evreux (Eure, France).

81

82 **Site presentation and reference survey**

83 The *Gisacum* religious center is located 7 km east of the capital city of *Aulerques*
84 *Ebuovices* (now Evreux in Normandy). There, inside a great hexagonal sacred polygon
85 covering 230 ha, stands a series of public buildings and temples, the dwelling and craft
86 activities being situated outside the polygon (Guyard and Lepert 1999). This site is located
87 over a great plateau where the superficial layer corresponds to flint clay. Above this clay, of
88 around 15 Ω .m resistivity, the archaeological remains have a variable thickness and may
89 surpass 100 Ω .m in resistivity (Aubry 2003). No remains are visible at the surface, and only
90 the thermal baths building and the main temple have been excavated.

91 The test took place in the area of the *fanum* (outlined blue rectangle in Figure 1) and to
92 the North over a plot named 'Terre noire'. The thickness of the archaeological layer in the
93 *fanum* remains around 90 cm. The test compares the results obtained using a three-depth
94 multipole array called ARP© (Automatic Resistivity Profiling), to those obtained with the
95 DualEM (see description in the next paragraph).

96 The ARP© is used generally for pedological and archaeological applications (Dabas,
97 2009); it weighs 400 kg and has 8 spike-wheels. The system is pulled with a quad-bike over a
98 grass cover and the measurement location is obtained using a combination of dGPS system
99 and a Doppler-radar system that allows checking the distance along the profile. The electrical
100 map obtained with the ARP© system is presented in Figure 1 where all buried structures are

101 clearly shown as resistive anomalies in black: walls of the Gallo-roman temple (*fanum*), walls
102 around the thermal complex (*peribola*), dwelling houses to the West delimited by roads (*via*),
103 aqueducts to the South bringing water to the thermal complex and to the temple of water
104 (*nymphaea*); see Dabas *et al.* (2005) for a complete explanation of archaeological features.

105

106 **Instrument and tests**

107 *Theoretical responses*

108 The DualEM-421S is a multi-receiver EMI instrument (DualEM sensor manual 2010)
109 operating at 9 kHz frequency. Its main characteristic is the provision of one horizontal
110 transmitter loop (TX coil, Figure 2) with three pairs of receivers. In each pair, the first
111 receiver is horizontal, allowing horizontal coplanar, HCP, and, by rotation of all the
112 apparatus, vertical coplanar, VCP, configuration measurements. The second receiver coil is
113 radial to the transmitter allowing perpendicular, PERP, configuration measurements. The first
114 pair is located at 1m and 1.1m from the transmitter, the second at 2 and 2.1 m and the third at
115 4 and 4.1 m. Measurements are triggered by time (up to 10 measurements by second) and
116 other data are available: temperature inside the tube, voltage of the battery and the two
117 rotation angles (roll and pitch).

118 Considering the usual resistivity range of soils and superficial formations, the
119 instrument belongs to the ‘Low Induction Number’ (LIN) EMI group (this number compares
120 the electromagnetic diffusion range characterized by the skin depth, $\sqrt{\frac{2}{\sigma\mu\omega}}$ where σ is the
121 conductivity, ω the angular frequency and μ the magnetic permeability, to the geometrical
122 dimensions of the considered problem: inter-coil spacing, depths of the targets, layer
123 thicknesses). For this group, the conductivity response is mainly in quadrature with the
124 primary field which facilitates the measurement of the magnetic susceptibility in phase. In
125 Figure 3, we illustrate the theoretical variations of the coil responses (defined as the ratio of

126 secondary field to primary field at the receiver coil, measured in parts per million) versus
127 conductivity or susceptibility for a homogeneous soil. In this figure and all along the text, the
128 responses are calculated without approximation using the complete mathematical expressions
129 of the secondary field (Thiesson *et al.* 2014). The left part shows the variation of the response
130 versus susceptibility (for a soil with a fixed 0.01 Sm^{-1} conductivity). The right part shows the
131 variation of the response when changing the conductivity (for a fixed susceptibility of $40 \cdot 10^{-5}$
132 SI in-phase and 0 in quadrature). The coil height is set to 0.1m above ground level and the
133 frequency to 9 kHz. It can be observed that all the dependences are quasi-linear: the
134 sensitivity of the phase responses to the magnetic susceptibility variations are good,
135 respectively, the sensitivity of the quadrature responses to the soil conductivity are also good.
136 But, as frequency, soil conductivity or coil separation increases the quadrature response does
137 not remain linear. For each transmitter-receiver separation, the conductivity limit above which
138 the discrepancy with a linear response surpasses a 10% threshold is illustrated in Table 1 for
139 both HCP and PERP geometries (for a 0.1 m height above a homogeneous ground). It must be
140 underlined that the HCP configuration is far more affected by this limit than the PERP one.
141 Consequently, it could be stated that, (except for 4m HCP pairs) in non-salted soils, the
142 dependences remain linear while this relationship is altered in saline environments (Beamish
143 2011). It must also be noted that for greater separations (4m), the amplitude of the responses
144 generated by the soil conductivity are significantly higher than those generated by the soil
145 susceptibility (Fig. 3): this would necessitate a very accurate phase separation of the
146 instrument responses.

147 The absolute calibration was controlled by measuring the DualEM response to a small
148 metallic conductive and non-magnetic sphere as a function to the height and distance to the
149 transmitter coil. Comparing the theoretical response of the sphere (not detailed in this paper,
150 see Thiesson *et al.* 2014) with the measured in-phase response (Figure 4a) and-b)), it appears

151 that most of the points exhibits an error less than $\pm 5\%$ except for the highest values where we
152 can hypothesize some non-linearity in the electronics.

153 Using a metallic sphere, no response should be observed in the three quadrature
154 components. Figure 4c and 4d show that the phase separation of the devices is better than
155 10% (except for the points in the dashed area but which are mainly located in the area of very
156 low amplitudes where noise is predominant). These two tests have shown that the device used
157 is reliable and that the measurements are accurate.

158 The 6 receivers correspond to six different depths of investigation that permit a 6
159 points 'geometrical sounding' at each measurement location. Consequently, the DualEM-421
160 can be used in a wide variety of near surface applications: soil, environmental, engineering
161 and archaeological studies. As with other Slingram instruments (Tabbagh 1986(c)), it also
162 allows detecting buried metallic objects. Several studies have already been published in
163 archaeology and soil studies (Simpson *et al.* 2009, Monteiro Santos *et al.* 2010, Saey *et al.*
164 2012, De Smedt *et al.* 2013) but none of them discussed in detail the ability of the instrument
165 for 2D/3D resistive features characterization.

166 The values displayed by the instrument are specific to the sampling volume of each
167 transmitter-receiver configuration, but the volumes are not the same for susceptibility as for
168 conductivity measurements. For the 6 in-phase channels, these values are directly expressed
169 as the ratio of the secondary field to the primary field in ppt (part per thousands). For the 6
170 quadrature channels, they are transformed to pseudo electrical conductivity by multiplying

171 each channel value by: $-\frac{4}{\mu_0 \omega L^2}$, where ω is the angular frequency, μ_0 the vacuum magnetic

172 permeability and L the separation between the coils corresponding to each channel. This
173 formula is valid only for coils at $h=0$, *i.e.* for a center of coils at ground surface. Even if the
174 instrument is laid on the ground, this hypothesis is not valid (the distance between the center

175 of the coils and the bottom of the boom is 4,5cm). When the instrument is towed on a cart, the
176 distance to the ground is 31.5cm.

177 Consequently:

178 (1) We do not use the pseudo electrical conductivities (output of DualEM), but the quadrature
179 magnetic field ratios using Table 2 (multiplication by a series of coefficients deduced from

180 $-\frac{4}{\mu_0 \omega L^2}$; also quoted in DualEM-421S User's manual, 2010),

181 (2) One must keep in mind that the difference between the output of DualEM and the
182 apparent conductivity is more important in HCP than in PERP and that this importance
183 increases with the average conductivity of the soil (clayey and even more salted soils). This is
184 a consequence of: i) greater non-linearity of the responses when the ground conductivity
185 increases, ii) the elevation of the instrument above ground level. For example, when
186 considering a 20 mS/m ground, the instrument at h=0.045 m height will measure in the HCP 4
187 m channel -5.340 ppt in quadrature and thus display 18.79 mS/m while at h=1 m (carried by
188 the operator) it will display 15.70 mS/m (-4.464 ppt). Under the same conditions, the PERP
189 4.1 m channel would deliver a -5.814 ppt at h=0.045m and display 19.5 mS/m while at h=1 m
190 it would measure -3.328 ppt and display 11.15 mS/m.

191 Therefore, it will be necessary to retranslate the displayed pseudo conductivity values
192 into field ratios when attempting to retrieve quantitative results (following Thiesson *et al.*
193 2014).

194

195 *Field stability tests*

196 A cart was designed, Figure 5, to pull the DualEM-421S with a quad bike on the field
197 at two possible clearances, 0.10 and 0.27 m from the bottom of the tube to the ground surface.
198 The distance between the quad-bike and the cart was tested so that no disturbance was
199 measured due to the presence of the quad-bike with engine switched off.

200 In order to measure the potential noise due to the engine of the quad-bike, DualEM
201 measurements were recorded continuously (10 Hz) during a few minutes with and without
202 engine working (Table 3). It can be observed that the engine has no effect and that the
203 standard deviation remains limited. PERP 1.1m configuration shows a much higher deviation
204 for an unknown reason. This experiment shows also that no “short time” drift (defined as a
205 variation within a 5 minutes time lapse, the time that would be needed to acquire for example
206 a 300m profile at 1m/s) was noticed. “Long term” drift (defined as the variations within the
207 the entire time scale of a survey, typically several hours) was not studied here and is the
208 subject of numerous papers ((Delefortrie *et al.* 2014, Dos Santos and Porsani 2011, Beamish,
209 2011). In archeology, we found that only “short time” drifts are really a problem because they
210 create artefacts in the maps that superpose to the short wavelength of archaeological
211 anomalies (typically several meters, that would correspond to a few seconds in the time
212 domain when using a speed of acquisition of the order of 1m/s). Long time drift, if ever
213 noticed, is easily corrected by standard image filtering procedures.

214 The roll is another important aspect when the instrument is pulled on uneven terrain (the pitch
215 angle also acquired with DualEM was not considered in this study). Figure 6 shows the
216 theoretical relative quadrature out of phase signal change versus roll (grey lines). The
217 maximum roll acceptable is 10 degrees (if we consider a maximum 2% deviation in the
218 quadrature out of phase signal) which defines the limit that must be respected in the field, but
219 PERP is more sensitive to the roll effect than HCP and there is a negative shift of the coil
220 orientation by reference to the vertical. This test also allows verifying the exact coil position
221 inside the housing tube. From the theoretical and practical point of view, the HCP
222 configuration is less sensitive to roll because, for an α roll angle, the HCP emission is reduced
223 by $\cos\alpha$ and the received field is also reduced by $\cos\alpha$ thus the HCP response is reduced by
224 $\cos^2\alpha$, but the rotation generated also a VCP response which, following the same reasoning, is

225 multiplied by $\sin^2\alpha$. As for small values of h , HCP and VCP quadrature responses are very
226 close (the lower the instrument, the closer their values), the roll effect is very limited. On the
227 contrary, in PERP configuration only the transmitter coil is tilted and the response is reduced
228 in $\cos\alpha$ (Figure 6). At least, as the spacing increases, it appears that the theoretical HCP
229 sensitivity to roll decreases, but practically appears another effect (probably due to some lack
230 in the stiffness between the boom and the sensor itself). Fortunately, the roll variation of the
231 signal with the PERP appears to be very repeatable and therefore could be corrected from the
232 value of roll measured.

233

234 *Comparison with vertical electrical soundings*

235 Six DC resistivity Vertical Electrical Soundings (VES) have been achieved to
236 recognize the vertical layering of the site (Figure 7). All exhibit the same distribution: a three
237 layer model, where, beneath a topsoil and an 'archaeological' layer of around 1m thickness
238 and [35-40 Ωm] resistivity, exists a conductive flint clay layer of thickness varying between 2
239 and 3 m and of resistivity between 10 and 20 $\Omega\text{.m}$ above a saturated chalk third layer of [30-
240 40 $\Omega\text{.m}$]. At the same locations, the conductivity responses of DualEM-421S were recorded
241 versus the height above ground surface. This test between 0.045m and 2.25m allowed, for
242 each geometrical configuration, both to assess the noise level (by the interquartile distance)
243 and to compare the displayed pseudo-conductivity values (transformed in ppt using the Table
244 1 coefficients) with the theoretical responses calculated from the VES interpreted parameters.
245 In Table 4 the comparison for the VES number 3 is detailed where $\rho_1=41.3 \Omega\text{m}$, $e_1=0.6 \text{ m}$,
246 $\rho_2=16.5 \Omega\text{.m}$, $e_2=2.06 \text{ m}$ and $\rho_3=26.9 \Omega\text{.m}$.

247 It can be observed in Table 4 that firstly the dispersion of the measurement is limited
248 to 1% of the response and independent of both the height above the ground surface and of the
249 receiver orientation. Secondly, the correlation with the theoretical data is very good for the

250 relative variations but exhibits different offsets for the different channels (if one hypothesizes
251 that the gain is the same for the different channels and is accurate). Thirdly, the comparison
252 with the approximate cumulative responses (McNeill 1980) confirms that this approximate
253 model is not relevant for quantitative interpretation.

254

255 *Reminder about the investigation depths*

256 We do not consider here the magnetic susceptibility (in-phase responses) changes
257 which are known to correspond to clearly lower depths of investigation than conductivity
258 changes (Scollar *et al.* 1990).

259 The quadrature measurement sensitivity of the HCP configuration to deep horizontal
260 layering is usually recognized as higher than that of the PERP configuration. However, the
261 comparison of their respective abilities is more complex for a 2D/3D structure: for example
262 the ability of the PERP configuration is generally higher than that of the HCP in the 3D case
263 (Tabbagh 1986(b)). Consequently we have decided not to use a definition which is linked to a
264 1D case and to approximate sensitivity functions using arbitrary levels (70% of the total
265 sensitivity for DOE, ‘depth of exploration’, and 50% for the DOI50, DualEM manual) even if
266 these latter permits very simple calculations. We think that the notion of investigation depth
267 we have chosen is of better use for surveyors as it is linked to the 3D case, to the full
268 calculation of the sensitivity functions, and to a standard level of ‘noise’ of 10%. Depth of
269 investigation is defined as to the depth where a change of more than 10% in apparent
270 conductivity can be measured when changing the target true electrical conductivity of the
271 target.

272

273 **DualEM results over Terre noire (wide mesh): spatial variations observed in the**
274 **superficial geologic context (1D inversion)**

275 In the DualEM–421S, the largest inter-coil separations allow a better characterization
276 of the superficial geological context (the first 5 m approximately) than presently available
277 mobile multi-pole resistivity systems (limited to approximately the first 2m). The 1D
278 inversion process has been applied with good results in a wide variety of near-surface studies
279 (Bendjoudi *et al.* 1998, Bobée *et al.* 2010). Even if it blurs the exact location of lateral limits
280 of the features (Guérin *et al.* 1996), it is very relevant for determining the thicknesses of the
281 different layers and their low frequency spatial variations, and in archaeological sites to
282 evaluate the bulk thickness of the remains.

283 The 1D inversion (program QwInv1D, UMR METIS, using an iterative least square
284 inversion (Guérin *et al.* 1996)) is here applied to the wide mesh (5m) EM data acquired in the
285 ‘Terre noire’ plot (Fig. 1) with nine different data series corresponding to HCP 1m, VCP 1m,
286 PERP 1.1m, HCP 2m, VCP 2m, PERP 2.1m, HCP 4m, VCP 4m and PERP 4.1m (all recorded
287 at 0.27m elevation and acquired continuously in two successive surveys). The original
288 pseudo-conductivity data were pre-processed to transform them into magnetic field ratio at a
289 height of 0.27m. Then, at each point of the surveyed area, the 9 experimental ratio values
290 were inverted to obtain the 5 following parameters: resistivity and thickness of the first layer,
291 resistivity and thickness of the second layer and resistivity of the third layer. This process was
292 done for the whole plot (3906 points for the 5 x 5 m² mesh). With a modal value of 6% for the
293 relative RMS error, the modal values of the first layer parameters are 41 Ω m and 1.0 m, those
294 of the second layer are 13 Ω m and 2.3 m and the resistivity of the third layer 39 Ω m. The
295 spatial variations of these parameters are presented in Figure 8. The first layer has a limited
296 resistivity which can be interpreted by an absence of stone remains except in the northern part
297 where such remains can be expected. A marked deep resistivity variation is present in the
298 second and third layer; it probably corresponds to a huge change in the chalk hardness that
299 outcrops with a reduction of the second layer thickness.

300 To take into account the fact that, with the DualEM 421S, VCP configuration would
301 necessitate a second survey, we repeat the inversion with 6 data only corresponding to HCP
302 and PERP configurations. The results obtained are quasi –identical to those of the previous 9
303 data inversion with a modal value of 5.5% for the relative RMS error, 41 Ωm and 1.02 m for
304 the first layer 12.3 Ωm and 2.3 m for the second layer and 40 Ωm for the third.

305

306 **DualEM results over the Fanum (fine mesh): detection of 3D resistive targets (3D**
307 **inversion and comparison with DC maps)**

308 An extensive survey was done by pulling the DualEM on its cart with a 0,1m clearance (only
309 in HCP configuration for the TX coils) over the *fanum*. Acquisition is in a continuous mode
310 (10Hz) following parallel lines separated by 0.5m and in zig-zag mode (10m between the
311 blocks of lines). Average velocity being 6.7 km.h⁻¹, the distance between acquisition points is
312 0.19 m. Positioning of data is done by a GNSS system using broadcasted differential
313 corrections from a base station situated less than 200m from the *fanum* area. The horizontal
314 relative accuracy is estimated around 1cm. The survey was done in 2 hours and 1 minute
315 (distance covered: 12780m; area 0,52ha). The statistics for the roll angle are such that no
316 correction was applied (mean=-3.8°, interquartile distance=7.7°).The change of temperature
317 was less than 2°C during the two hours survey (median value =24.1°C). Positioning of the
318 data was done knowing the distance between the GPS and the different coils (Figure 2) and
319 also the time lag between DualEM and GPS. Data, converted to apparent electrical resistivity
320 for comparison with DC maps, were re-interpolated over a square mesh of 0.2 x 0.2m using a
321 bi-cubic spline function. We have decided not to use any 1D or 2D filters for these data in
322 order to show the raw data.

323

324 *Qualitative comparison of EM and DC maps*

325 In Figure 9 are presented the 6 apparent resistivity maps obtained with the DualEM,
326 together with the two apparent resistivity maps obtained with ARP©. As stated before, no
327 filtering was undertaken to generate these maps. No outlier points exist for the 4m coils. The
328 first outliers appear for the HCP2m coil and for both 1m coils. HCP configuration seems to be
329 more sensitive to these shallow outliers. We think that most of these outliers come from small
330 metallic objects spread all over the surface (the in-phase maps, not shown in this article, show
331 high amplitude anomalies at the same position; this area is very close to an air-force base that
332 was bombed during last world war). These outliers can be easily filtered by a 2D median
333 filtering. Some “long term” drift does appear (see the ‘bands’ in PRP 4.1m for example) in
334 these maps but does not affect the visibility of archaeological features in this case.

335 It can be observed that:

- 336 (1) Except for the larger coil separation in EMI, all the main resistive features are clearly
337 delineated even if the apparent resistivity anomaly magnitudes are more limited in
338 EMI than in DC.
- 339 (2) The apparent resistivity values are in good coherence with the depth of investigation
340 expected from sensor geometries. The 1D point by point inversion applied in that area
341 with a three layer model delivers the following mode values: 70 Ωm and 0.2 m for the
342 first topsoil layer, 26 Ωm and 1.3 m for the ‘archaeological’ layer and 20 Ωm for the
343 flint clay.
- 344 (3) As expected from both theoretical (Tabbagh 1986(b)) and physical (Frischknecht *et al.*
345 1991) models, the HCP responses are oscillating and exhibit two parallel resistivity
346 maxima beside each wall and a stronger minimum just above it. This ‘triple arch’
347 response would be confusing when direct reading of the map in absence of 2D/3D
348 modelling reference, but it does not impede the interpretation (see below).

349 (4) In agreement with the expected depths of the walls (tops at around 0.3m), the ‘best’
350 results are given by PERP 1m configuration.

351

352 *3D interpretation of the wall of the Cella: quantitative comparison of inverted 3D DC data*
353 *and DualEM data*

354 To go further in the comparison, one can apply over a limited area, the two-step
355 1D/3D rapid inversion described in (Brinon *et al.* 2012). This process is based for 1D on an
356 analytical forward modelling and for 3D on the moment method. The latter is relevant to
357 quantify the responses of simple-shape features in both EM modelling, for resistivity and
358 magnetic susceptibility contrasts (Tabbagh 1985) and in DC resistivity modelling (Dabas *et*
359 *al.* 1991). This approach allows a direct comparison of the resistivity contrasts of the detected
360 features.

361 We consider, for example, a small area limited by X= 519018.56 and 519023.36 in
362 easting and Y= 145232.66 and 145236.26 in northing, delineated in Figure 9, where can be
363 observed the external *cella* wall. Using the data of the three channels of the ARP© one
364 obtains for the surrounding terrain 1D model $\rho_1=70 \Omega\text{m}$, $e_1=0.2 \text{ m}$, $\rho_2=32 \Omega\text{m}$, $e_2=1.3 \text{ m}$ and
365 $\rho_3=19 \Omega\text{m}$. For the wall, the inversion gives an orientation at 100° from X axis, a section of
366 $1.0 \times .88 \text{ m}^2$ for a center at X=519020.96, Y=145234.46 and Z=0.7m and a resistivity of 70.5
367 Ωm . The contrast ratio between the wall and the surrounding is thus 2.2.

368 Adopting these values, it is possible to calculate using the moment method the
369 responses that would be obtained for the different DualEM coil configurations. Due to the
370 depth of the wall, the calculations are limited to HCP 1m, HCP 2m, PERP 1.1 m and PERP
371 2.1m. The results of the comparison with the DualEM measured profile are presented in
372 Figure 10 (Hs/Hp values in ppt) for the four coil configurations. The agreement is good both

373 for magnitude and width. It also confirms the anomaly shapes with a single maximum in
374 PERP configuration and the three arch responses for HCP configuration.

375 In this example, it is clear that the walls are detected by the four coil configurations as
376 resistive features, and would be correctly interpreted using only one of each. This result
377 confirms what has been already established (Thiesson *et al.* 2009) for a smaller instrument in
378 VCP configuration. In fact, the relative variation of amplitude of the responses above the wall
379 is rather limited, 10% for the PERP and about 7% for HCP compared to 25% observed in the
380 ARP© channels. However, here again, the relevant criterion is the signal to noise ratio, i.e. the
381 relative sensitivity of instruments to the searched-for features as opposed to the unwanted
382 features ('geophysical noise'). In the Vieil-Evreux site where the global resistivity is of the
383 order of 100Ωm and lower, the signal to noise ratio can be comparable to that of the DC
384 resistivity method.

385

386 **Conclusions**

387 We presented here a test of a single frequency domain EMI instrument over an
388 archaeological site in order to better grasp the capabilities of this exploration technique. Even
389 though this technique has been used for fifty years in archaeological survey and is sensitive to
390 both magnetic susceptibility and electrical conductivity, it has been, and still is, less applied
391 than magnetic and earth-resistance prospection.

392 The Vieil-Evreux roman site remains (mainly walls) are situated in a sedimentary (clay)
393 geological context where the conductivity response is significantly high; consequently the two
394 major limitations of EMI cannot be considered in the present study but they merit to be
395 recalled in this conclusion because they, at least for a part, explain the restricted use of the
396 technique. They are: (1) metallic objects and features disturb measurements and practically
397 disqualify its application specially in urban contexts and (2) in highly resistive soil contexts

398 the conductivity response is too small and difficult to separate from the quadrature magnetic
399 susceptibility in the quadrature response (Tabbagh 1986(c)) in spite of recent research works
400 aiming to overcome this problem (Simon *et al.* 2015).

401 The series of tests conducted with the DualEM in Vieil-Evreux have allowed verifying that
402 the electronic and external electromagnetic noises are fairly rejected, the phase of the received
403 signals are well defined, the absolute calibration is fine, the drifts limited and the roll effect
404 characterized. Measured quadrature data are in full accordance with the theoretical complete
405 calculations taking the true resistivities and layer thicknesses as input data (from separate
406 VES). Using this instrument, it is thus possible to assess the advantage and drawbacks of the
407 method itself within the particular objective of the detection of stone built resistive features.

408 The first point to be emphasized is that frequency domain multi-receiver EMI instruments,
409 lighter and easier to move than multi-pole DC, can be used to characterize the superficial
410 geological context with a sufficient depth of investigation. This aspect is important for
411 extended archaeological sites as it allows a first rough estimation of the volume of the
412 archaeological remains. The second point is the ability to detect and characterize walls. They
413 are detected with 1m and 2m coil separations and the corresponding apparent resistivity maps
414 clearly show the plan of the fanum at Vieil-Evreux. The corresponding signal variations are
415 lower than with DC resistivity but as the geophysical noise is also lower, the observed signal
416 to noise ratios are roughly equivalent. Nevertheless, referring to 3D theoretical modelling is
417 required to correctly interpret the HCP raw data.

418 Finally the Vieil-Evreux test verifies that with sufficient fine meshing and a close referencing
419 to modelling, stone built resistive features can be surveyed with frequency domain multi-
420 receiver EMI.

421

422

423

424 **Acknowledgements**

425 *We wish first to thank MADE (Mission Archéologique du Département de l'Eure) for*
426 *accessing this test site, GEOCARTA for giving free access to the instrument; to Rick Taylor*
427 *and Scott Holladay for their corrections of this manuscript and to the reviewers for their*
428 *helpful contributions.*

429

430 **References**

- 431 Aubry L., 2003. Acquisition, traitement et restitution des données d'une reconnaissance
432 archéologique : la ville gallo-romaine de Vieil-Evreux. Thèse Université Pierre et Marie
433 Curie.
- 434 Beamish D., 2011. Low induction number, ground conductivity meters: a correction
435 procedure in absence of magnetic effects. *Journal of applied Geophysics*, 75, 244-253.
- 436 Bendjoudi H., Weng P., Guérin R., Pastre J.-F., 2002, Riparian wetland of middle reach of the
437 Seine river (France): historical development, investigation and present hydrologic
438 functioning. A case study. *Journal of Hydrology*, 263, 131-155.
- 439 Bobée C., Schmutz M., Camerlynck C., Robain H., 2010. An integrated geophysical study of
440 the western part of the Rochechouart-Chassenon impact structure, Charente, France. *Near*
441 *Surface Geophysics*. 8-4, 259-270.
- 442 Bonsall J., Fry R., Gaffney C., Armit I., Beck A., Gaffney V., 2013. Assessment of the CMD
443 Mini-Explorer, a new low-frequency Multi-coil Electromagnetic device, for Archaeological
444 Investigations, *Archaeological Prospection*, 20, 3, 219-231.
- 445 Bossuet G., Thivet M., Trillaud S., Marmet E., Laplaige C., Dabas M.; Hulin G., Favard A.,
446 Combe L., Barres E., Lacaze S., Aubry L ; Chassang M., Mourot A., Camerlynck C., 2012.
447 City map of ancient Epomanduodurum (Mandeure-Mathay, Franche-Comté, Eastern France):
448 contribution of geophysical prospecting techniques. *Archaeological Prospection*, **19**, 261-280.
- 449 Brinon C., F.-X. Simon, A. Tabbagh, 2012, Rapid 1D/3D inversion of shallow resistivity
450 multipole data: examples in archaeological prospection. *Geophysics*, **77-3**, E193-E201.
- 451 Dabas M., 2009. Theory and practice of the new fast electrical imaging system ARP©, in:
452 *Seeing the Unseen, Geophysics and Landscape Archaeology*, Campana and Piro eds., CRC
453 Press, Taylor and Francis Group, 2009, 105-126.

454 Dabas M., Guyard L. and Lepert T. 2005. Gisacum revisité : croisement géophysique et
455 archéologie. In « Géophysique et archéologie », *Dossiers de l'Archéologie* **308**, 52-61.

456 Dabas M., Tabbagh A., 2003. A comparison of EMI and DC methods used in soil mapping-
457 theoretical considerations for precision agriculture, p. 121-129, in: *Precision Agriculture*, ed.
458 J. Stafford and A. Werner, Wageningen Academic Publishers, Muencheberg, 783p.

459 Dabas, M., Tabbagh A., Tabbagh J. 1994. 3D inversion in subsurface electrical surveying, I:
460 Theory. *Geophysical Journal International*, **119**, 975–990.

461 Delefortrie, -S., -De -Smedt, -P., -Saey, -T., van -De -Vijver, -E., van -Meirvenne, -M., 2014.
462 An efficient calibration procedure for correction of drift in EMI survey data. *Journal of*
463 *Applied Geophysics*, 110, 115-125.

464 Dos Santos, -V.R.N., -Porsani, -J.L., -2011. -Comparing -performance -of -instrumental
465 drift correction -by linear -and -quadratic -adjusting -in inductive -electromagnetic- data.
466 *Journal -of Applied Geophysics*, 73, 1-7.

467 Elwaseif, M., and L. Slater, 2010, Quantifying tomb geometries in resistivity images using
468 watershed algorithms. *Journal of Archaeological Science*, **37**, 1424-1436.

469 De Smedt P., Saey T., Lehouck A., Stichelbaut B., Meerschman E., Islam M. M., van De
470 Vijver E, van Meirvenne M. 2013. Exploring the potential of multi-receiver EMI survey for
471 geoarchaeological prospection: a 90 ha dataset. *Geoderma*, **199**, 30-36.

472 DualEM-421S User's manual, 2010, Milton, Ontario, Canada, 32p.

473 Flageul S, Dabas M, Thiesson J, Réjiba F, Tabbagh A. 2013. First in situ tests of a new
474 electrostatic resistivity meter. *Near Surface Geophysics*, **11-3**: 265-273.

475 Frischknecht F. C., Labson V. F., Spies B. R., Anderson W. L. 1991. Chapter 3 Profiling
476 methods using small sources. *Electromagnetic methods in applied geophysics, vol. 2,*
477 *Application, part A* 105-252. Edited by M. N. Nabighian, SEG, Tulsa.

478 Gaffney C.F., Gater J.A., Linford P., Gaffney V.L., White R. 2000. Large-scale systematic
479 fluxgate gradiometry at the roman city of Wroxeter. *Archaeological Prospection*, **7**: 81-99.

480 Gebbers R., Lück E., Dabas M., Domsch H., 2009. Comparison of instruments for
481 geoelectrical soil mapping at the field scale, *Near Surface Geophysics*, **7**, 179-190.

482 Guérin R., Méhéni Y., Rakontodrasoa G., Tabbagh A., 1996. Interpretation of Slingram
483 conductivity mapping in near surface geophysics: using a single parameter fitting with 1D
484 model. *Geophysical Prospecting*, **44-2**, 233-249.

485 Guyard, L, Lepert T., 1999. Le Vieil- Evreux, ville sanctuaire gallo-romaine, *Archeologia*
486 **359**, 20-29.

487 McNeill J. D. 1980. Electromagnetic terrain conductivity measurement at low induction
488 numbers. *Geonics ldt technical note* TN-6, 15p.

489 Monteiro-Santos F. A., Triantafilis J. Bruzgulis K., 2011. A spatially constrained 1D
490 inversion algorithm for quasi-3D conductivity imaging: Application to DUALEM-421 data
491 collected in a riverine plain. *Geophysics*, **76-2**, B43-B53.

492 Novo A., Dabas M., Morelli G., 2012: Fast High-Resolution Archaeological Mapping:
493 STREAM X tested at Vieil-Evreux (France), *Archaeological prospection*, **19**, 3, 179-189.

494 Papadopoulos N. G., Tsokas G. N., Dabas M., Yi M-J., Kim J-H., Tsourlos P., 2009. Three-
495 dimensional inversion of automatic resistivity profiling data. *Archaeological Prospection*, **16**,
496 267-278.

497 Saey T., van Meirvenne M., De Smedt P., Neubauer W., Trinks I., Verhoeven G., Seren S.,
498 2013. Integrating multi-receiver electromagnetic induction measurements into the
499 interpretation of the soil landscape around the school of gladiators at Carnuntum. *European*
500 *Journal of Soil Science*, **64**, 716-727.

501 Saey T., De Smedt P., Meerschman E., Islam M.-M., Meeuws F., van De Vijver E., Lehouck
502 A., van Meirvenne M., 2012. Electrical conductivity depth modelling with a multi-receiver
503 EMI sensor for prospecting archaeological features. *Archaeological Prospection*, **19**, 21-30.

504 Scollar I, Tabbagh A., Hesse A. Herzog I., 1990, Archaeological prospection and remote
505 sensing. *Cambridge University Press*, 674p.

506 Simon F.-X., Sarris A., Thiesson J., Tabbagh A., 2015. Mapping of quadrature magnetic
507 susceptibility/magnetic viscosity of soils by using multi-frequency EMI. *Journal of Applied*
508 *Geophysics*, *120*, 36-47.

509 Simpson D., van Meirvenne M., Saey T., Vermeersch H., Bourgeois J., Lehouck A., Cockx
510 L., Vitharana U. W 2009. Evaluating the multiple coil configurations of the EM38DD and
511 Dualem-21S sensors to detect archaeological anomalies. *Archaeological Prospection*, **16**, 91-
512 102.

513 Tabbagh A. 1985. The response of a tree dimensional magnetic and conductive body in
514 shallow depth E.M. prospecting. *Geophysical. Journal of the Royal Astronomical. Society*, 81-
515 1, 215-230.

516 Tabbagh A. 1986(a). Sur la comparaison entre la prospection électrique et trois méthodes de
517 prospection électromagnétique pour la détection de contrastes de résistivité associés aux
518 structures archéologiques. *Prospezioni Archeologiche*, **10**: 49-63.

519 Tabbagh A., 1986(b). What is the best coil orientation in the slingram electromagnetic
520 prospecting method? *Archaeometry*, **28-2**, 185-196.

521 Tabbagh A., 1986(c). Applications and advantages of the Slingram electromagnetic method
522 for archaeological prospecting. *Geophysics*, 51-3, 576-584.

523 Thiesson J, Dabas M, Flageul S 2009. Detection of resistive features using towed slingram
524 electromagnetic induction instruments. *Archaeological Prospection*, **16**: 103-109.

525 Thiesson J., Kessouri P., Schamper C., Tabbagh A. 2014. Calibration of frequency-domain
526 electromagnetic devices used in near-surface surveying. *Near Surface Geophysics*, **12**, 481-
527 491.
528

529 **Figure captions**

530 Figure 1: Composite apparent resistivity map of Vieil-Evreux Gallo-Roman site (RM15-
531 MuCEP and ARP®channel2); ARP® system in the upper right with array configuration;
532 studied area outlined in blue.

533

534 Figure 2: General scheme of coil configuration in the DualEM-421S instrument (cm units).

535

536 Figure 3: Theoretical in-phase (continuous lines) and quadrature responses (dotted lines)
537 versus magnetic susceptibility (left side) and electrical conductivity (right side) for the
538 different HCP and PERP distances from the transmitter in the DualEM-421S.

539

540 Figure 4: a) Results of the calibration with the sphere at $z=60$ cm above the device, b) Results
541 of the calibration with the sphere at $z=23$ cm above the device
542 c) Phase separation for the Perp geometries d) Phase separation for the HCP geometries

543

544 Figure 5: DUALEM-421S on its cart pulled by a quad.

545

546 Figure 6: Relative quadrature out of phase signal changes (theoretical: grey line, practical
547 dark diamonds) as function of roll angle (in degree) a) HCP 4m, b) PERP 4.1m, c) HCP 2m,
548 d) PERP 2.1m, e) HCP 1m, f) PERP 1.1m.

549

550 Figure 7: Position of the 6 VESs and example of inversion for VES N°3

551

552 Figure 8: Terre noire plot: spatial variations of the 5 parameters inverted from the 9
553 quadrature data.

554

555 Figure 9: Apparent resistivity maps over the *fanum* for the 6 different DualEM-421S
556 quadrature channels and two of the ARP© channels (3D inversion area outlined in blue).

557

558 Figure 10: Comparison between the experimental (bold line) values obtained over the small
559 test area and the responses calculated using the result of the ARP© data inversion (thin line).

560

561

562 **Table captions**

563 Table 1: Conductivity values above which the quadrature response differs with more than
564 10% from linear response.

565

566 Table 2: Coefficients allowing the computation of magnetic field ratio.

567

568 Table 3: Measurement dispersion (resistivity in Ωm) at a fixed point with (a) and without (b)
569 engine-turned on.

570

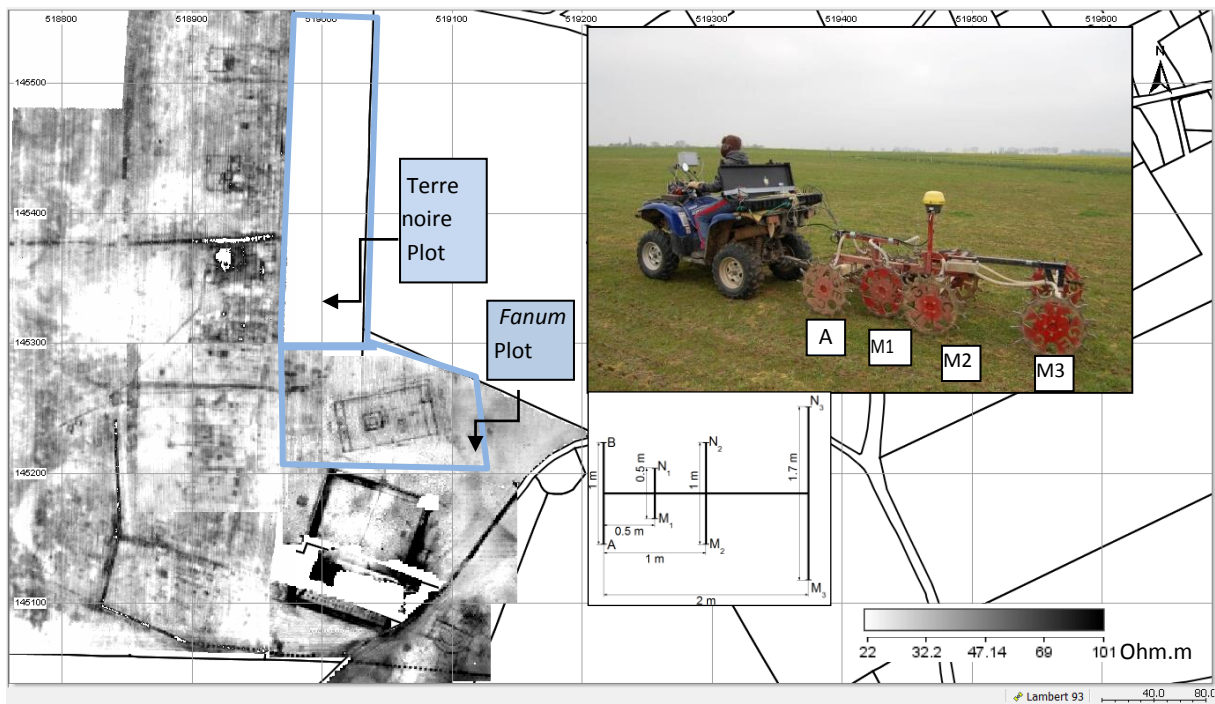
571 Table 4: Experimental dispersion and median values observed versus instrument height above
572 the ground surface at VES 3 location and comparison with the values deduced from the VES
573 three layer interpretation where $\rho_1=41.3\ \Omega\text{m}$, $e_1=0.6\ \text{m}$, $\rho_2=16.5\ \Omega\text{m}$, $e_2=2.06\ \text{m}$ and $\rho_3=26.9$
574 Ωm using the complete expression of the secondary field (Thiesson *et al.* 2014) and using the
575 approximate cumulative sensitivity response (Wait 1962, Mc Neill 1980). The order of the
576 coil geometries is the following: HCP 1 m, PERP 1.1 m, HCP 2 m, PERP 2.1 m, HCP 4 m
577 and PERP 4.1 m.

578

579

580

581



582

583

584 Fig. 1

585

586

587

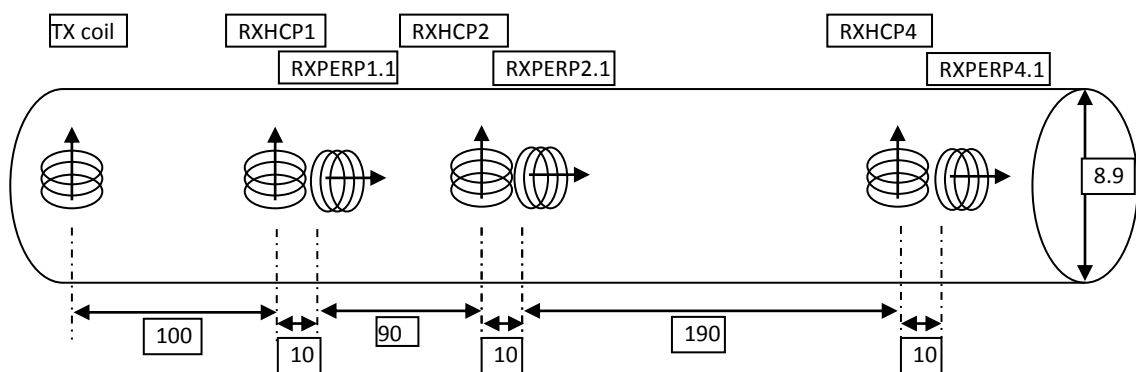
588

589

590

591

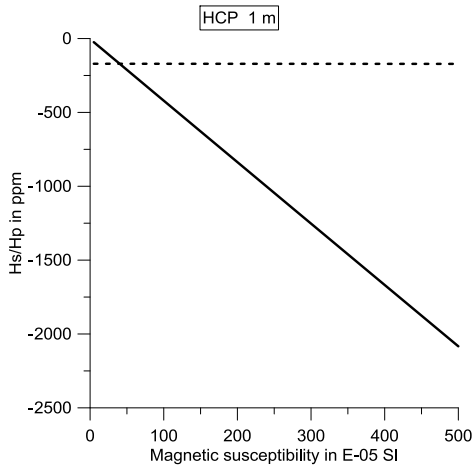
592



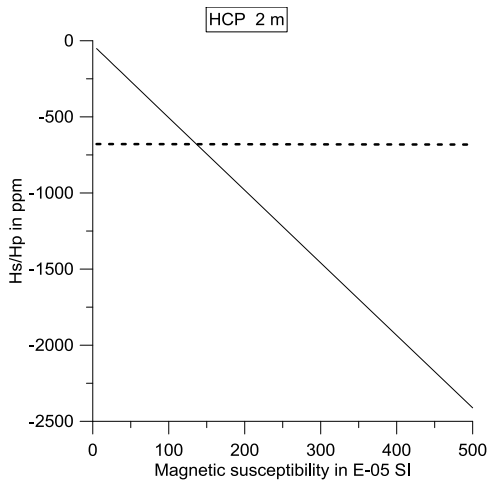
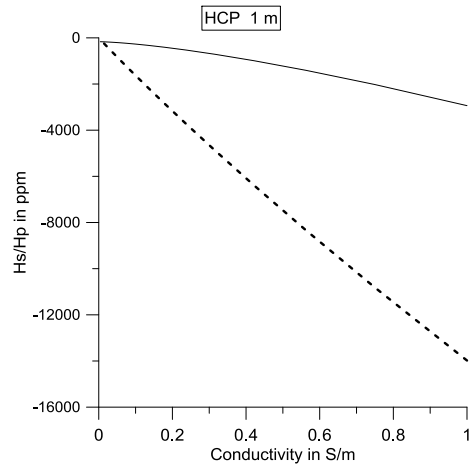
593 Fig. 2

594

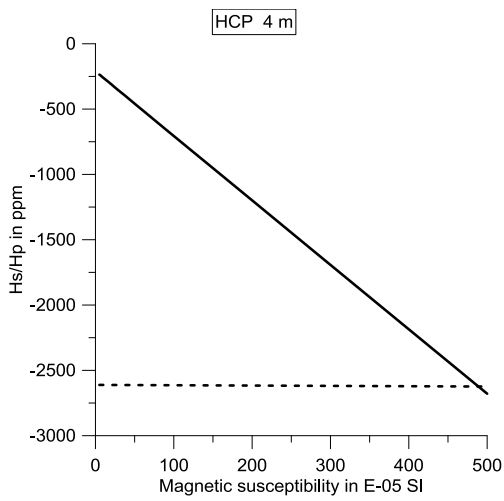
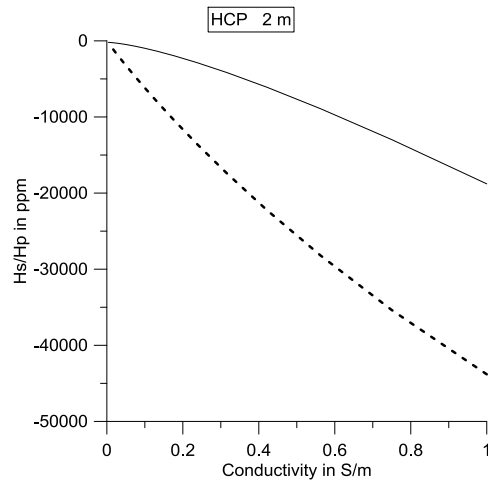
595



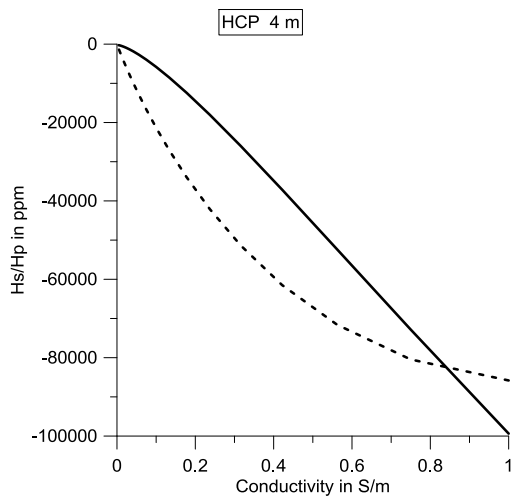
596

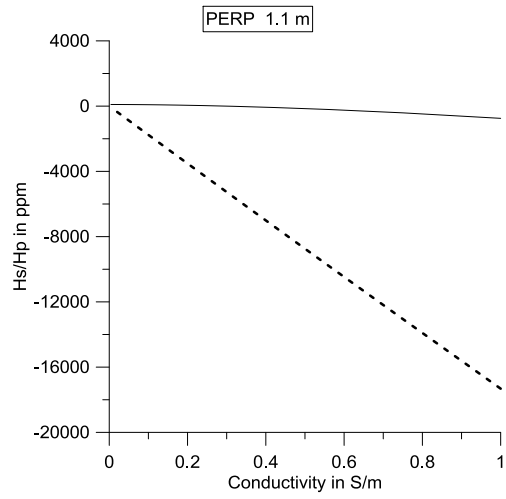
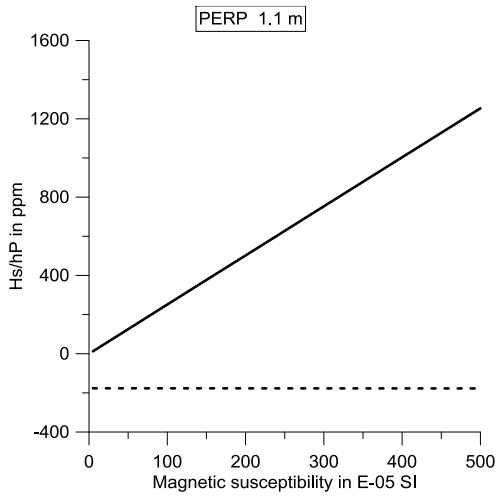


597

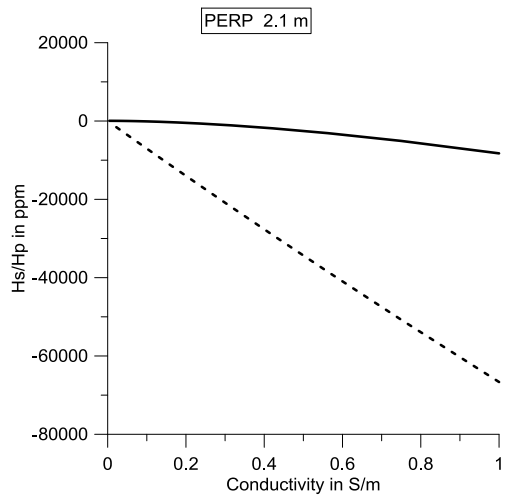
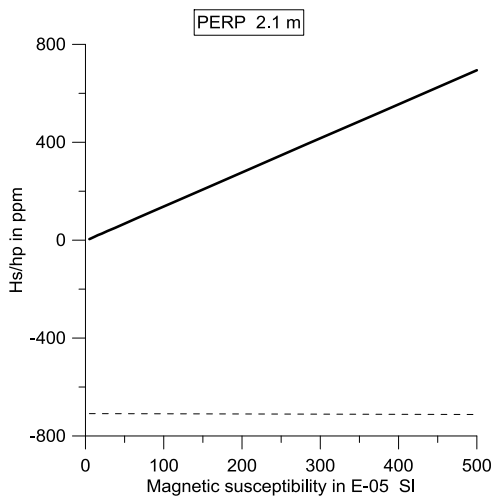


598

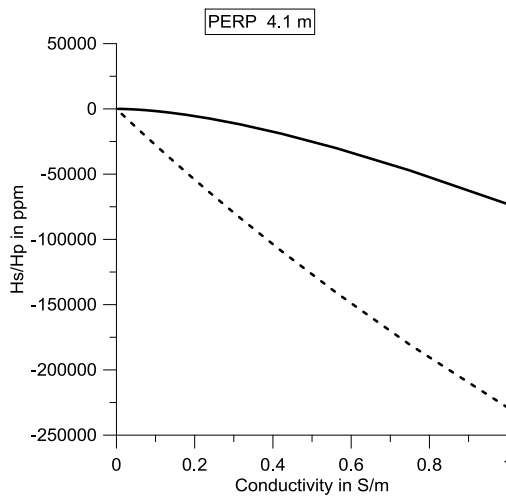
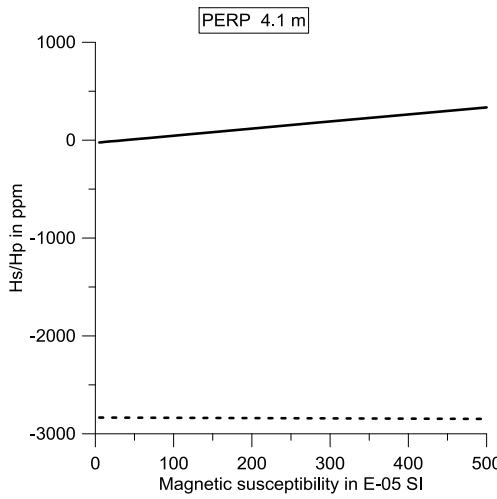




599



600

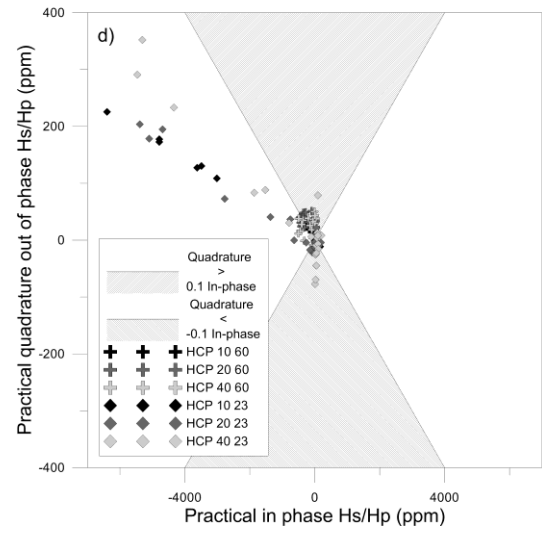
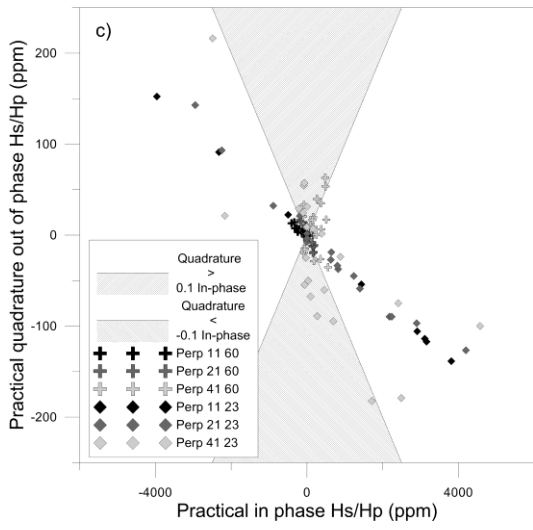
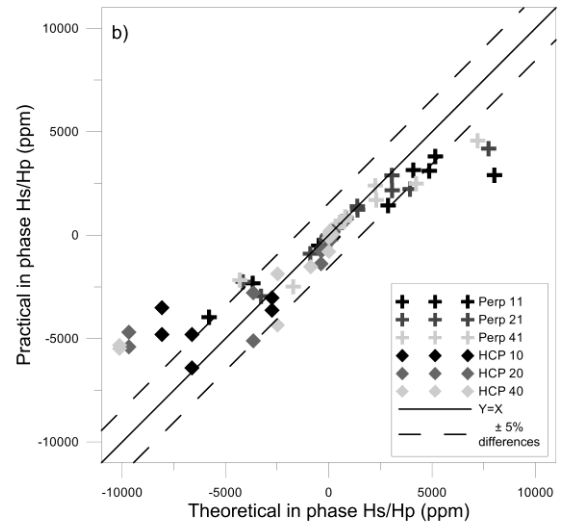
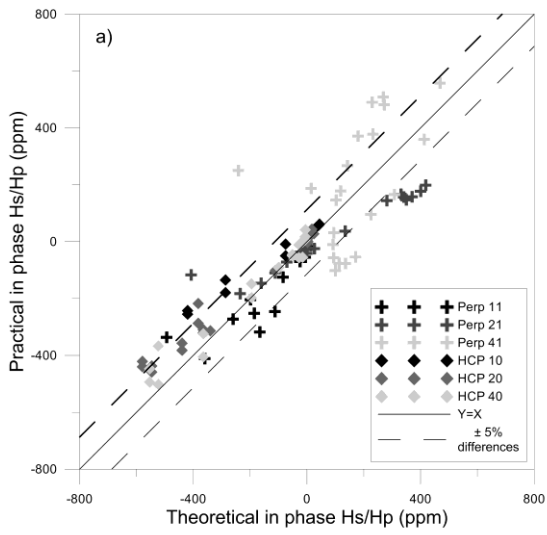


601

602 Fig. 3

603

604



605

606

607 Fig. 4

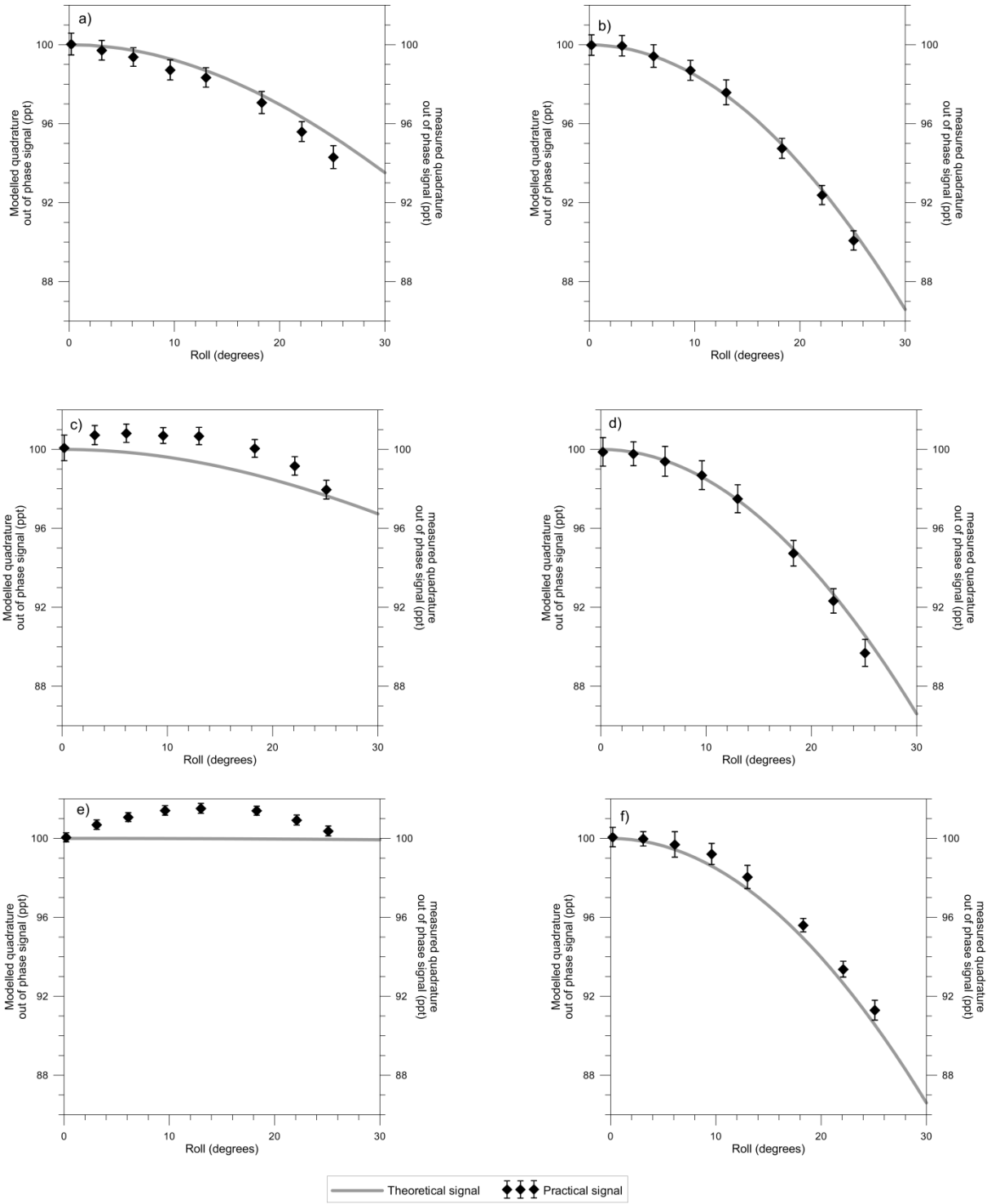
608



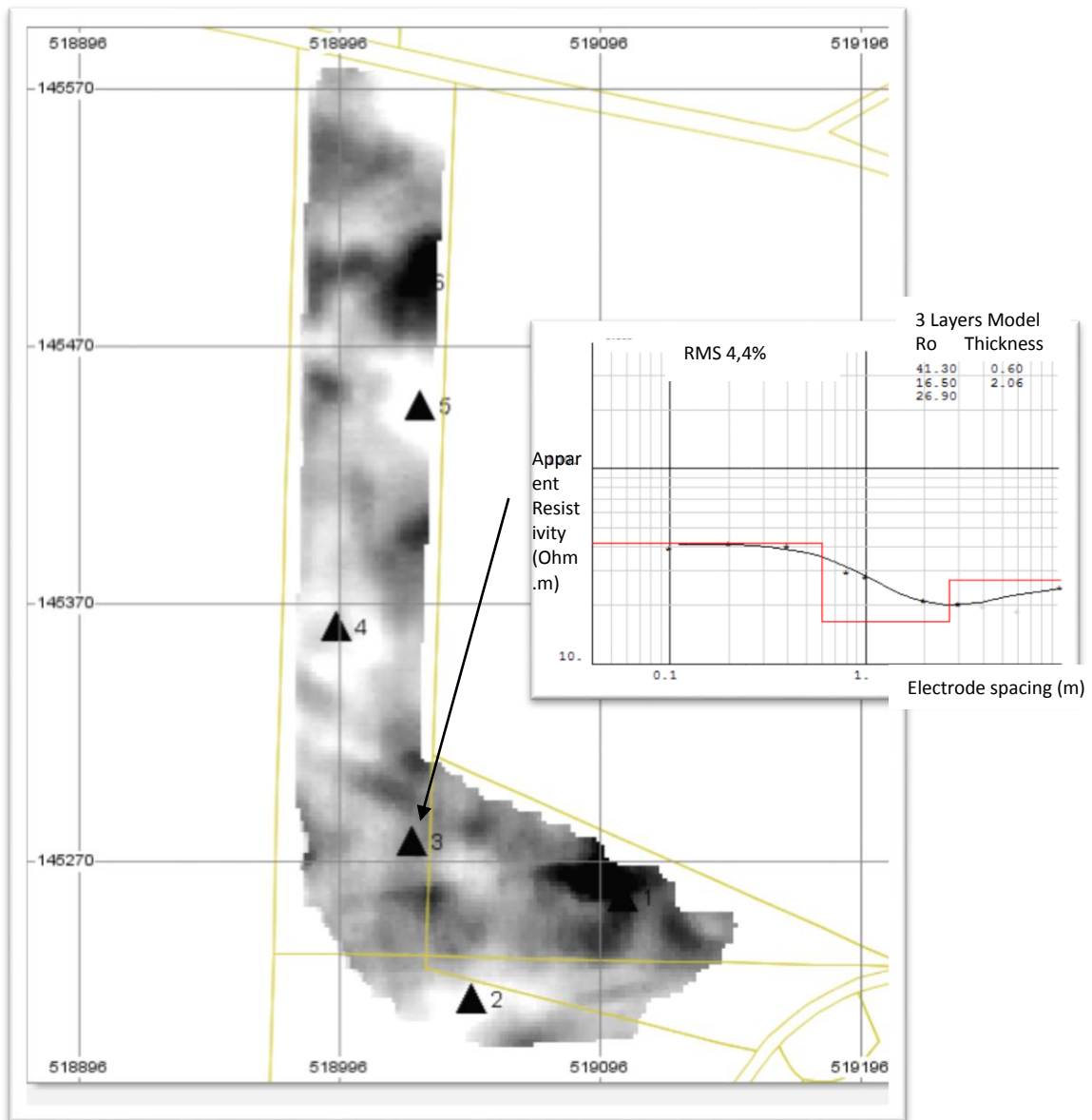
609

610 Fig. 5

611



616

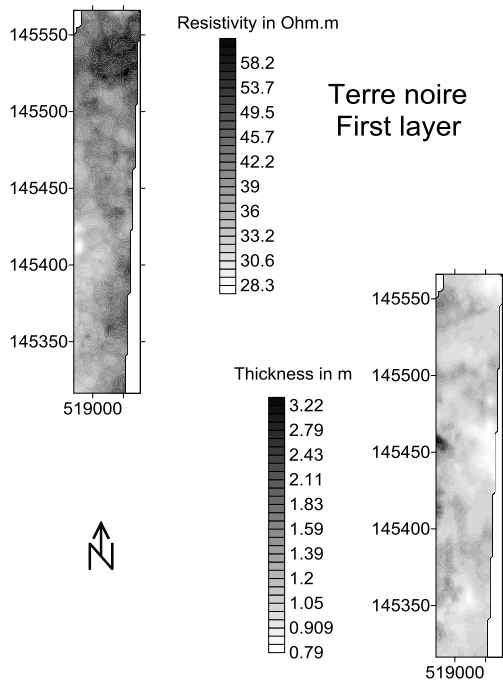


617

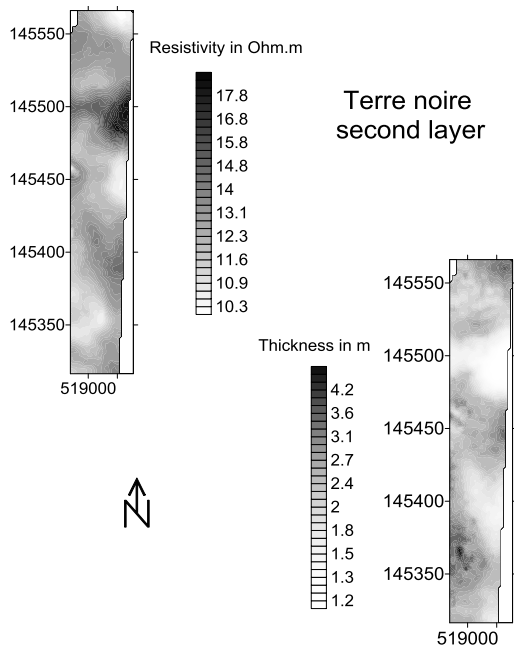
618 Fig.7

619

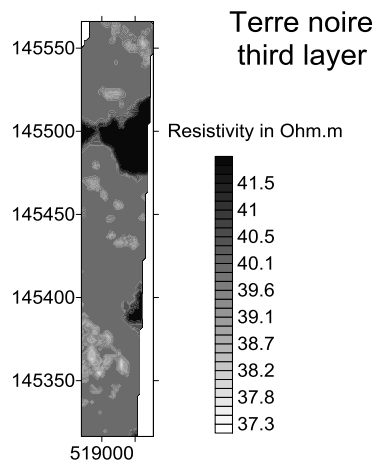
620



621



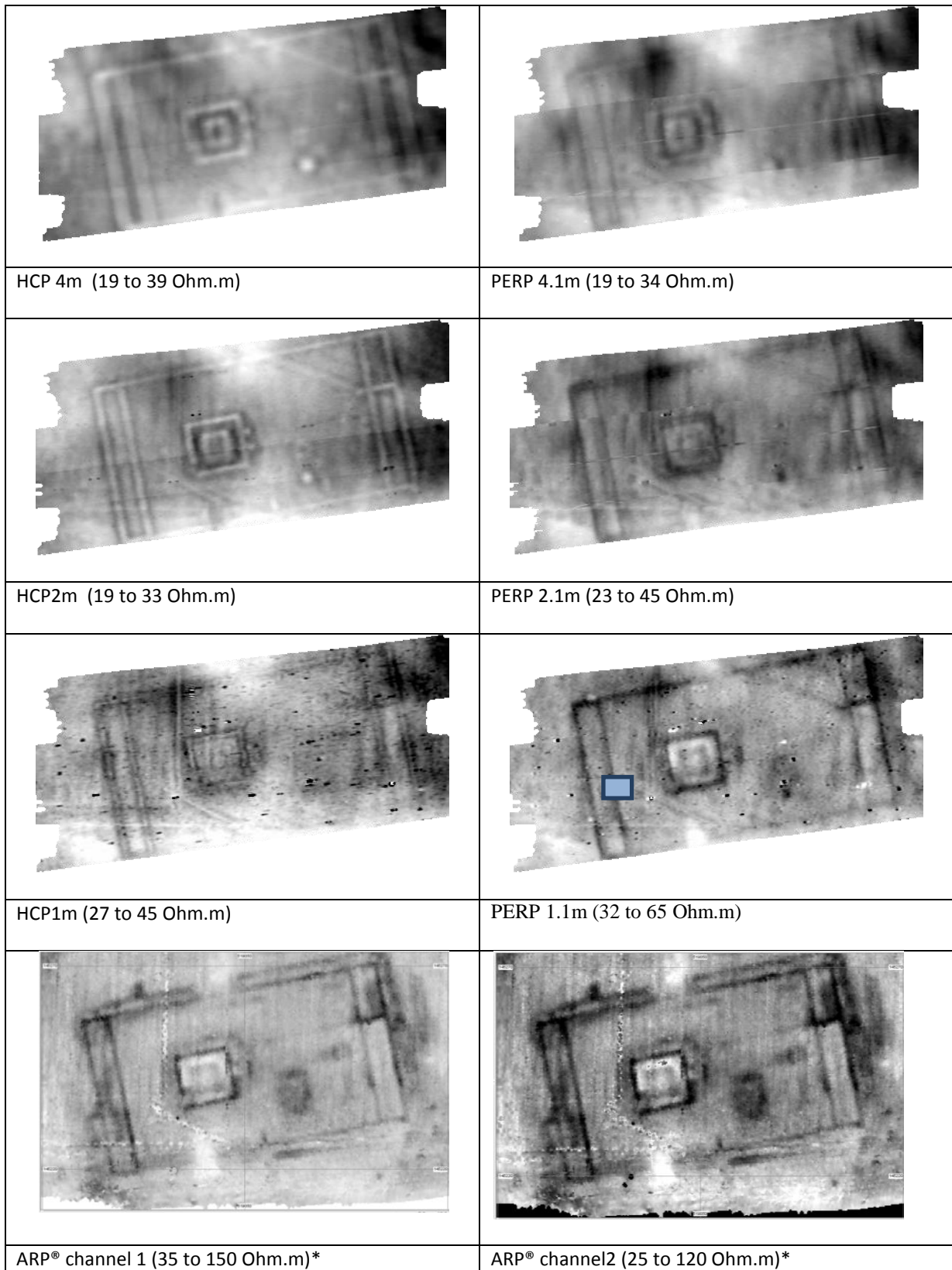
622



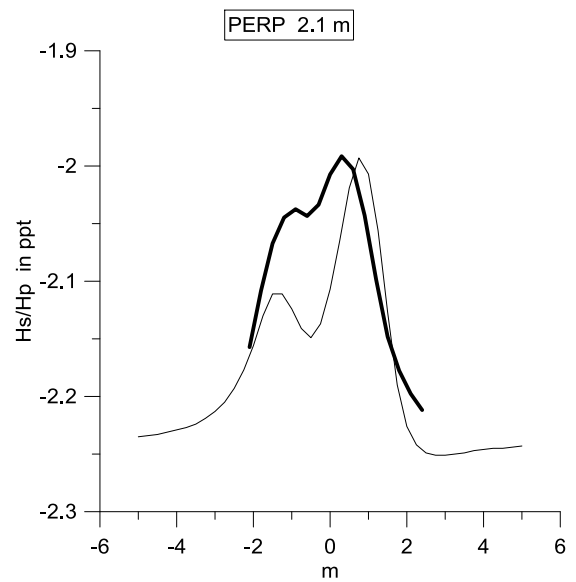
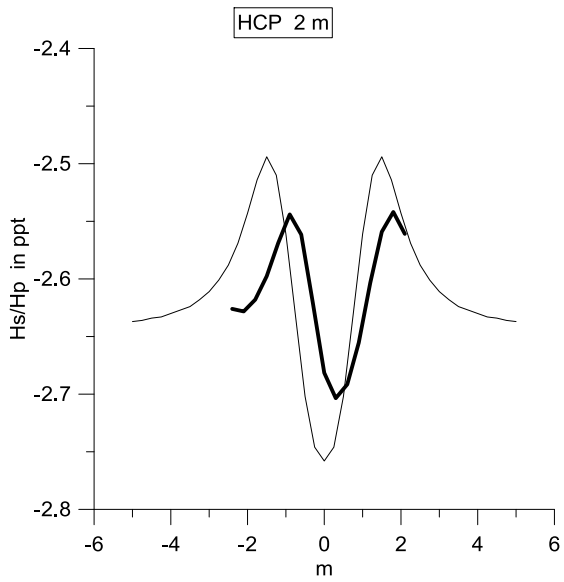
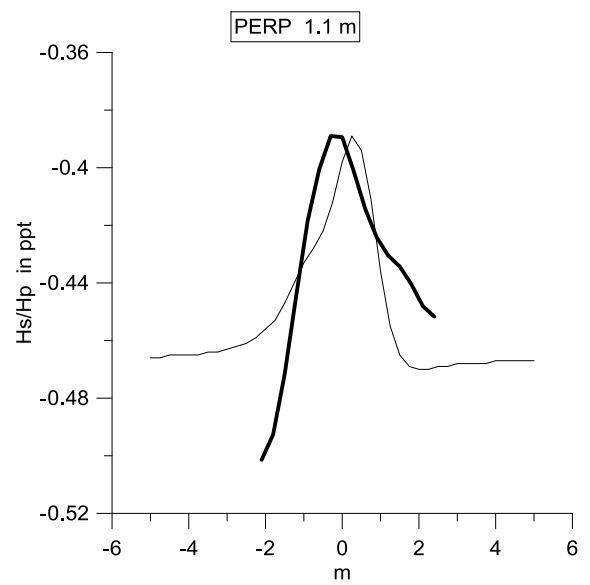
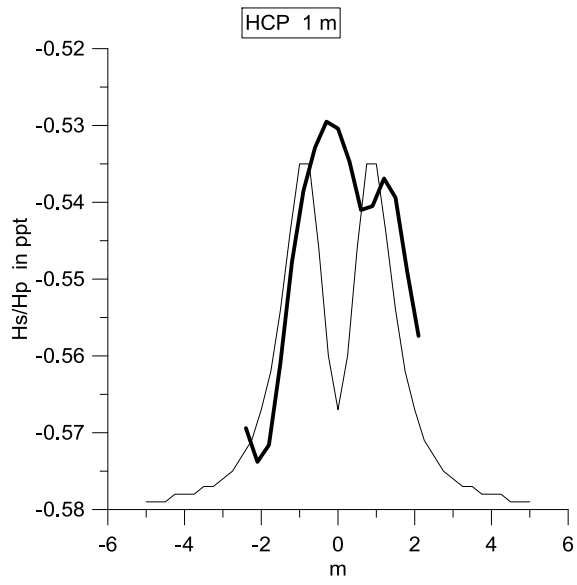
623

624 Fig. 8

625



627 Fig. 9



628

629 Fig. 10

630

Max	29.1	27.2	28.7	82.6	43.1	333.3
Min	27.6	26	26.2	74.1	39.5	243.9
Average	28.32	26.60	27.53	78.02	41.28	279.7
Variance	0.029	0.033	0.023	1.842	0.150	95.9
Standard deviation	0.17	0.18	0.15	1.36	0.39	9.75

637 (b)

638

639 Table 4

Height above ground surface (m)	Experimental Dispersion Third quartile – first quartile (ppt)	Median experimental value (ppt)	theoretical value deduced from VES data (ppt)	Approximate cumulative responses from VES data (ppt)
0.045	0.005	0.552	0.722	1.169
	0.004	0.516	0.654	0.796
	0.028	2.936	3.098	4.125
	0.024	2.670	3.043	3.667
	0.057	11.246	11.043	10.769
	0.119	12.661	13.151	17.062
0.112	0.005	0.531	0.691	1.143
	0.004	0.449	0.570	0.757
	0.028	2.922	3.036	4.100
	0.031	2.466	2.820	3.577
	0.057	11.161	11.003	10.921
	0.119	12.243	12.662	16.849
0.179	0.005	0.518	0.655	1.113
	0.004	0.400	0.493	0.714
	0.028	2.887	2.966	4.067
	0.031	2.286	2.609	3.480
	0.057	11.076	10.948	11.054
	0.090	11.705	12.180	16.628
0.246	0.005	0.493	0.617	1.081
	0.004	0.355	0.425	0.668
	0.028	2.823	2.889	4.026
	0.023	2.122	2.410	3.378
	0.057	10.962	10.878	11.168
	0.090	11.227	11.707	16.397
0.347	0.004	0.452	0.560	1.031
	0.002	0.284	0.340	0.598
	0.021	2.709	2.765	3.952
	0.031	1.864	2.134	3.214
	0.057	10.792	10.747	11.305
	0.119	10.541	11.016	16.034
0.449	0.005	0.409	0.506	0.983
	0.004	0.224	0.273	0.531
	0.009	2.567	2.633	3.868

	0.031	1.644	1.884	3.045
	0.057	10.565	10.587	11.406
	0.090	9.824	10.343	15.651
0.75	0.007	0.304	0.377	0.872
	0.004	0.120	0.152	0.365
	0.028	2.211	2.239	3.602
	0.031	1.167	1.307	2.546
	0.057	9.883	9.990	11.524
	0.090	8.062	8.538	14.456
1.05	0.005	0.231	0.291	0.801
	0.004	0.073	0.094	0.249
	0.028	1.962	1.889	3.360
	0.023	0.830	0.925	2.099
	0.085	9.088	9.280	11.451
	0.060	6.659	7.020	13.225
1.35	0.005	0.182	0.232	0.756
	0.004	0.049	0.063	0.167
	0.028	1.585	1.600	3.165
	0.031	0.564	0.674	1.723
	0.057	8.293	8.531	11.273
	0.090	5.449	5.772	12.015
1.65	0.004	0.146	0.192	0.725
	0.004	0.032	.044	0.107
	0.028	1.322	1.367	3.013
	0.031	0.446	0.506	1.415
	0.057	7.526	7.795	11.049
	0.090	4.539	4.963	10.869
1.95	0.005	0.117	0.162	0.704
	0.004	0.024	0.033	0.062
	0.028	1.145	1.181	2.898
	0.023	0.353	0.390	1.163
	0.057	6.816	7.103	10.817
	0.090	3.792	3.954	9.810
2.25	0.005	0.098	0.139	0.690
	0.004	0.17	0.026	0.026
	0.028	1.003	1.030	2.810
	0.031	0.290	0.307	0.955
	0.057	7.696	6.467	10.597
	0.090	3.195	3.506	8.845

640

641

642